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# Nitrous oxide dominates greenhouse gas emissions from hydropower's reservoirs in China from 2020 to 2060



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#### ABSTRACT

China is ambitious to increase its hydropower share to mitigate climate changes. The greenhouse gas (GHG) emissions from hydroelectric reservoirs may hinder the climate goal. The spatio-temporal patterns of such emissions under future climate changes at the national scale are not clearly addressed. In this study, we evaluate these emissions from 79 hydroelectric reservoirs across China (61.22 % of the national hydropower generation) in 2020, covering carbon dioxide (CO<sub>2</sub>), methane (CH<sub>4</sub>), and nitrous oxide (N<sub>2</sub>O), using the G-res (Greenhouse Gas Reservoir) tool and Integrated Model to Assess the Global Environment–Dynamic Global Nutrient Model (IMAGE-DGNM). A random forest (RF) model is also used to project the emissions in the period of 2020 to 2060 under Shared Socioeconomic Pathway (SSP) scenarios. The results indicate that the carbon intensity (CI) and areal flux varied largely. The reservoirs located in low-altitude areas and older reservoirs generation is a key factor influencing the N<sub>2</sub>O emissions. The projection shows that these emissions will increase by 1.30 %, 6.63 %, and 17.33 % in 2060 compared to 2020 under the SSP119, SSP245, and SSP585 scenarios, respectively, in which CH<sub>4</sub> has the largest growth. Finally, implications toward reduction in such emissions are discussed.

#### Nomenclature

IHA	International Hydropower Association	YRB	Yangtze River Basin	
SLRB	Songhua and Liaohe River Basin	RF	Random forest	
CO2	Carbon dioxide	SwB	Southwest Basin	
HuRB	Huaihe River Basin	YeRB	Yellow River	
			Basin	
N2O	Nitrous oxide	PRB	Pearl River Basin	
IPCC	Intergovernmental Panel on Climate	CH4	Methane	
	Change			
CI	Carbon intensity	HRB	Haihe River	
			Basin	
GRanD	Global reservoir and dam database	SeB	Southeast Basin	
SSPs	Shared Socioeconomic Pathways	GHG	Greenhouse gas	
IEA	International Energy Agency	CB	Continental	
			Basin	
IMAGE- DGNM	Integrated Model to Assess the Global Environment-Dynamic Global Nutrient	WRT	Water residence time	
	Model			

 

 (continued)

 GWP
 Global Warming Potential
 G-res
 Greenhouse Gas Reservoir

 DLEM
 Dynamic Land Ecosystem Model
 ResME
 Reservoir Methane Emissions

 News
 Nutrient Export from WaterShed
 Vertice

#### 1. Introduction

To meet the increasing energy demand, hydropower has been identified as one of the climate-friendly sources (Liu et al., 2022). China has a unique advantage to promote its hydropower, owing to the huge theoretical hydropower potential and technically exploitable potential of 694 GW (15 % of the global total) and 542 GW, respectively (Sun et al., 2019). To meet the China's ambitious goal of achieving carbon neutrality in 2060, it has been proposed to enhance the share of hydropower to 20 % of the total energy production in 2050 to reduce the fossil fuel consumption (Liu et al., 2022). However, hydropower's

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Received 15 July 2024; Received in revised form 19 February 2025; Accepted 28 February 2025 Available online 1 March 2025 0043-1354/© 2025 Elsevier Ltd. All rights are reserved, including those for text and data mining, AI training, and similar technologies. reservoirs, with relatively large storage areas and long operational life compared to other types of reservoirs, result in stable and significant sources of greenhouse gases (GHGs) such as  $CH_4$ ,  $CO_2$ , and  $N_2O$  (Fearnside et al., 2012; Harrison et al., 2009; Hendzel et al., 2005; Liu et al., 2011; Rudd et al., 1993), which can hinder the climate mitigation goal. Therefore, evaluation and projection of the spatio-temporal patterns and related drivers of these emissions at the China's national scale can provide further insights toward the climate goal.

Previous estimates indicated annual emissions of 328 Tg CO<sub>2</sub>, 22 Tg CH<sub>4</sub>, and 0.71 Tg N<sub>2</sub>O from global reservoirs (Harrison et al., 2021; Wang et al., 2023), which represent approximately 1.3 % of the anthropogenic CO<sub>2</sub> equivalent emissions globally, which is comparable to those from biomass burning or rice paddies (Li and He, 2022). This highlights the importance for reduction in such emissions. Measurement-based and model-based approaches are commonly employed methods to evaluate these emissions. Floating chamber (Li et al., 2015; Teodoru et al., 2012) and gas chromatograph (Shi et al., 2020) are widely used measurement tools. Previous studies indicated that the emissions evaluated by these measurement-based approaches are significantly lower than the actual emissions as the GHG emissions from bottom water are generally not measured (Adams, 2005; Cheng et al., 2019). Model-based studies can provide more systematic and comprehensive estimations owing to the ability to consider multiple and region-specific factors, such as the hydrological and climatic conditions of the reservoir. Among the various models, such as the generalized linear model (Scherer and Pfister, 2016), Reservoir Methane Emissions (ResME) models (Delwiche et al., 2022), and HydroCalculator Tool (Vilela and Reid, 2017), the Greenhouse Gas Reservoir (G-res) model, based on empirical measurements, is recommend by Intergovernmental Panel on Climate Change (IPCC) for Tier 3 estimates. Levasseur et al. (2021) demonstrated that the G-res model is more appropriate to evaluate CO2 and CH4 from hydropower reservoirs. However, reservoirs can emit considerable levels of N<sub>2</sub>O (Wang et al., 2023), with a considerably higher global warming potential (GWP), which is not considered in the G-res model. Besides, previous models, including Dynamic Land Ecosystem Model (DLEM) (Li et al., 2024), mechanistic mass balance model (Maavara et al., 2019), and Nutrient Export from WaterShed (News) model (Kroeze et al., 2010), used to calculate N<sub>2</sub>O emissions, overlooked emissions from reservoir sediments, while Integrated Model to Assess the Global Environment-Dynamic Global Nutrient Model (IMAGE-DGNM) addresses this gap (Vilmin et al., 2020). In this regard, the integration of both G-res model and IMAGE-DGNM can provide a more holistic information regarding the GHG from reservoirs.

The GHG fluxes from the reservoirs are mainly determined by the related size, age, latitude, operating season, and organic matter of the reservoir (Barros et al., 2011; Beaulieu et al., 2014; Soumis et al., 2004; Tremblay et al., 2004). In this context, the future emissions can be projected considering the significant heterogeneities of these drivers. Soued et al. (2022) used the G-res model to project the cumulative emissions from global reservoirs and reported that the CH4 bubbling and degassing flux will dominate in the reservoir-induced radiative forcing in the future. Li and He (2022) showed that the GHG emissions of reservoirs can increase approximately 0.11 Pg  $CO_2$  eq yr<sup>-1</sup> when the dams under construction and planning are considered. However, the climate changes can significantly impact the GHG emissions from reservoirs. For instance, the drawdown of water levels in high-altitude reservoirs, which are located in climate-sensitive area, may increase CH<sub>4</sub> and N<sub>2</sub>O emissions (Zhang et al., 2024). Higher temperatures can accelerate the decomposition of organic matter in soils and siltation in reservoirs, increasing the GHG emissions (Yang et al., 2014). An increased precipitation may lead to a higher degree of inundation, promoting anaerobic decomposition of organic matter to produce more GHG emissions (Rosa et al., 2004). The previous studies seldom address the future climate scenario, which may limit the understanding on climate impacts for these emissions. To this end, the Shared Socioeconomic Pathways (SSPs), proposed by IPCC, have been intensively employed to

model the future emissions under climate changes.

In summary, the studies on GHG emissions from hydropower reservoirs generally ignored the  $N_2O$  emission and impacts of climate change on these emissions. To fill this gap, we estimate the emissions of CH<sub>4</sub>, CO<sub>2</sub>, and N<sub>2</sub>O from 79 hydroelectric reservoirs across China, based on the G-res tool and IMAGE-DGNM. A sensitivity analysis is conducted to ensure reliability of our results. We project the GHG emission of these reservoirs in the period of 2020 to 2060 based on the SSP scenario by using a random forest (RF) model. In addition, we discuss mitigation pathways.

#### 2. Methods

#### 2.1. Study reservoirs

We selected study samples that are regionally representative with available data. The 79 large hydropower stations have an installed capacity exceeding 500 MW, generate 61.22 % of the nation's annual hydropower output, and account for over 52 % of the total national installed hydropower capacity (NEA, 2016). The locations of the studied reservoirs include the China's main basins (Fig. 1): 49 in the Yangtze River Basin (YRB), 4 in the Songhua and Liaohe River Basin (SLRB), 10 in the Southwest Basin (SwB), 7 in the Yellow River Basin (YeRB), 6 in the Pearl River Basin (PRB), and 3 in the Southeast Basin (SeB). Among them, 42 hydropower stations are located at low altitudes (<1000 m), while the others are at high altitudes. The parameters of these hydropower stations, such as the annual energy output, reservoir capacity, and reservoir area, varied largely (Table S1), which can provide more holistic and region-specific results related to their GHG emissions.

#### 2.2. Data sources

The reservoir surface area, reservoir capacity, and geographic locations of existing reservoirs were derived from H. Zhang et al. (2023). Data for the hydropower stations (installed capacity, catchment area, dam height, and power generation) were obtained from National Energy Administration. The average depth, maximum depth, and nitrate content of the reservoir were obtained from previous studies (Guo et al., 2020; Hu et al., 2019; Ji et al., 2022; C. Li et al., 2019; Yue et al., 2017; Zhang et al., 2021; Zhao et al., 2024, 2019; Zhou et al., 2016). The water qualities (water temperature, total phosphorus, dissolved oxygen, ammonia nitrogen, total nitrogen) of the reservoirs were obtained from China National Environmental Monitoring Center. The river length was obtained from China Water Statistical Yearbook (2020). The soil carbon content was obtained from World Soil Information. The cumulative global horizontal radiance was obtained from Geographic Information System Remote Sensing. Monthly meteorological data in the period of 2020 to 2060 and streamflow were obtained from National Tibetan Plateau Data Center. As presented in Table S2, all data used to evaluate the GHG emission in 2020, except the concentration of NO<sub>3</sub> and dissolved N<sub>2</sub>O, are primarily from the year 2020. Owing to the lack of these two data, we sourced them from the literature for study years closest to 2020.

#### 2.3. Methods for calculation of the GHG emissions from reservoirs

#### 2.3.1. Evaluation of the $CO_2$ and $CH_4$ emissions from reservoirs

The  $CO_2$  and  $CH_4$  emissions from reservoirs were evaluated by the Gres tool, developed by the International Hydropower Association (IHA) and recommended by IPCC. The G-res tool evaluates the  $CO_2$  diffusive emissions,  $CH_4$  diffusive emissions,  $CH_4$  bubbling emissions, and  $CH_4$ degassing emissions (Prairie et al., 2021). Table 1 shows the factors used to calculate the GHG flux. The equations used in the G-res model have been reported by Prairie et al. (2021).



Fig. 1. The location of studied hydropower reservoirs in China.

 Table 1

 The input factors for calculating the GHG flux.

GHG flux	Factors
CO <sub>2</sub> diffusive emissions	River area before impoundment (%), effective temperature CO <sub>2</sub> ( °C), reservoir area (km <sup>2</sup> ), reservoir soil carbon content (kg C m <sup>-2</sup> ), total phosphorus in the reservoir ( $\mu$ g l <sup>-1</sup> )
CH <sub>4</sub> diffusive emissions	Littoral area (%), effective temperature $\rm CH_4$ ( $^\circ\rm C)$
CH <sub>4</sub> bubbling emissions	Littoral area (%), reservoir cumulative global horizontal radiance (kg C m <sup>-2</sup> period <sup>-1</sup> )
CH <sub>4</sub> degassing emissions	Reservoir area (km <sup>2</sup> ), annual runoff (mm yr <sup>-1</sup> ), catchment area (km <sup>2</sup> ), CH <sub>4</sub> diffusive emissions (g $CO_2 m^{-2} yr^{-1}$ ), water residence time (yr)

#### 2.3.2. Evaluation of $N_2O$ emissions from reservoirs

The N<sub>2</sub>O emission calculation for the reservoirs in this study was based on the IMAGE-DGNM model. It has been widely used in simulations of N flows across landscapes and waterscapes (Beusen et al., 2015; Liu et al., 2018). It can simulate biogeochemical cycling of multiple elements and forms in global watersheds by using the speciation of different nutrient sources. IMAGE-DGNM evaluates the nitrification in water ( $NIT_{N_2O}$ ), denitrification in water ( $DENIT_{N_2O}$ ), incomplete nitrifier denitrification ( $NIT\_DENIT_{N_2O}$ ), nitrification in sediments ( $BNIT_{N_2O}$ ), denitrification in sediments ( $BDENIT_{N_2O}$ ), and N<sub>2</sub>O exchange at the water–atmosphere interface ( $EXCH_{N_2O}$ ); detailed calculations consider the supplementary material. For each waterbody type within a grid cell, changes in the amount of inland water N<sub>2</sub>O (dN<sub>2</sub>O) during the time step (dt) can be expressed in a mass-balanced manner:

$$\frac{dN_2O}{dt} = L_{N_2O} - QC_{N_2O} + NIT_{N_2O} + NIT_{DENITN_2O} + DENIT_{N_2O} + BNIT_{N_2O} + BDENIT_{N_2O} + EXCH_{N_2O},$$
(1)

where  $L_{N_2O}$  is the N<sub>2</sub>O input to the waterbody (from headwaters, upstream grids, and external groundwater input N<sub>2</sub>O and atmospheric deposition input N<sub>2</sub>O), Q is the discharge (m<sup>3</sup>·h<sup>-1</sup>), and  $C_{N_2O}$  is the inland water dissolved N<sub>2</sub>O concentration (mol· $m^{-3}$ ). The net N<sub>2</sub>O production includes nitrification(*NIT*<sub>N<sub>2</sub>O), nitrifier</sub>

denitrification( $NIT_DENIT_{N_2O}$ ), and denitrification( $DENIT_{N_2O}$ ), which can be positive in the case of net N<sub>2</sub>O production, intermediate during an incomplete denitrification, or negative in the case of net N<sub>2</sub>O consumption via complete denitrification in the water column and that from nitrification( $BNIT_{N_2O}$ ) and denitrification in the bed sediment ( $BDENIT_{N_2O}$ ).  $EXCH_{N_2O}$  is the N<sub>2</sub>O exchange at the water–atmosphere interface.

In this study, GHG emissions were converted into carbon footprint based on the GWP. The GWP of N<sub>2</sub>O is 298 times the CO<sub>2</sub> equivalent, significantly higher than that of CH<sub>4</sub> (34 times the CO<sub>2</sub> equivalent) in a 100-year time horizon (Liu et al., 2019). The GHG emission intensity of the reservoir is measured by the areal flux rate and carbon intensity (CI) (Almeida et al., 2019; Harrison et al., 2021). The areal flux rate indicates the GHG emission per surface area over a given period (g CO<sub>2</sub> eq m<sup>-2</sup> yr<sup>-1</sup>), while the CI is used to evaluate the total GHG emission per generated MWh (kg CO<sub>2</sub> eq MWh<sup>-1</sup>).

#### 2.4. Projection of GHG emissions from hydropower reservoirs in 2060

The prediction in this study can be classified into the following steps, as shown in Fig. S1. First, both historical data related to air temperature, precipitation, river flow, water quality parameters, reservoir parameters, and sediment parameters from previous studies and predicted data related to air temperature and precipitation based on the SSP scenario were collected. Second, based on the RF model (detailed information considers the supplementary material), 80 % of the dataset related to temperature, precipitation, and river flow was selected as a training set, while the remaining 20 % of the dataset was selected as a test set for model training and validation. Third, the reservoir water temperature and reservoir area were projected according to previous studies (McMahon et al., 2007; Toffolon and Piccolroaz, 2015; H. Zhang et al., 2023). Finally, the G-res model and IMAGE-DGNM were employed to predict the GHG emissions from the reservoir.

GHG emissions from reservoirs may impact China's carbon neutrality target, which necessitates forecasts up to the year 2060. By predicting emissions to 2060, policymakers and stakeholders can prepare more systematic strategies to achieve the carbon neutrality goal. Our climate scenarios essentially cover a range of future climates (Chapman et al., 2022; Wang et al., 2024; Wu et al., 2024), including SSP119 (lowest vulnerability, lowest mitigation pressure, and lowest anthropogenic radiation forcing), SSP245 (moderate social vulnerability and moderate anthropogenic radiative forcing), and SSP585 (highest artificial radiation forcing), representing contrasting projections of future climate changes.

In this study, we did not consider future changes in power generation mainly owing to the following reasons. The hydropower generation is influenced by fluctuating electricity supply and demand driven by the gross domestic product (GDP) (Qin et al., 2020; Zhou et al., 2020) and electricity prices (Fan et al., 2020); the prediction of these factors can lead to a significant uncertainty, which is beyond the scope of this study. Besides, the accurate projection of hydropower generation under future climate conditions requires detailed monthly historical data on temperature, rainfall, and electricity generation, which are currently unavailable.

#### 3. Results

#### 3.1. GHG emissions from reservoirs in 2020

Fig. 2a presents the CO<sub>2</sub>, CH<sub>4</sub>, and N<sub>2</sub>O areal fluxes from the reservoirs. The studied reservoirs had a wide range of GHG area flux, from 297.18 to 8517.99 g CO<sub>2</sub> eq m<sup>-2</sup> yr<sup>-1</sup> (average  $\pm$  standard deviation = 1900.86  $\pm$  1766.40 g CO<sub>2</sub> eq m<sup>-2</sup> yr<sup>-1</sup>) in 2020. N<sub>2</sub>O dominated in the GHG emission (81.10 % in terms of CO<sub>2</sub> eq), followed by CO<sub>2</sub> (13.20 %) and CH<sub>4</sub> (5.70 %). The reservoirs of Yantan Plant (8517.99 g CO<sub>2</sub> eq m<sup>-2</sup> yr<sup>-1</sup>), Tianshengqiao 1st Plant (7362.62 g CO<sub>2</sub> eq m<sup>-2</sup> yr<sup>-1</sup>), and Changheba Plant (7219.38 g CO<sub>2</sub> eq m<sup>-2</sup> yr<sup>-1</sup>) were the top three emitters.

Fig. 2b presents the CIs of the reservoirs. The values of the studied reservoirs are 0.006 to 1296.97 kg CO<sub>2</sub> eq MWh<sup>-1</sup> (average  $\pm$  standard deviation = 60.19  $\pm$  176.21 kg CO<sub>2</sub> eq MWh<sup>-1</sup>). The top three hydropower reservoirs with the highest CIs are in Fengman reservoir (1296.97

kg  $CO_2$  eq MWh<sup>-1</sup>), Danjiangkou reservoir (632.1 kg  $CO_2$  eq MWh<sup>-1</sup>), and Xinanjiang reservoir (475.1 kg  $CO_2$  eq MWh<sup>-1</sup>). These reservoirs are older than 50 years. Older reservoirs tend to have higher CIs, primarily due to the decrease in the depth of the reservoir by the accumulation of sediment over time. The reservoir depth is a crucial factor affecting this indicator (Li and He, 2022). Additionally, the increasing age of hydropower stations may negatively impact the power generation efficiency, and further result in an increase in reservoir CI.

At the catchment scale, PRB had the highest GHG areal flux of 4884.24 g CO<sub>2</sub> eq m<sup>-2</sup> yr<sup>-1</sup> (Fig. 3a), while YeRB had the lowest, 1588.40 g CO<sub>2</sub> eq m<sup>-2</sup> yr<sup>-1</sup>. According to the categories of GHG, YRB had the highest CO<sub>2</sub> areal flux of 524.53 g CO<sub>2</sub> eq m<sup>-2</sup> yr<sup>-1</sup>, while YeRB had the lowest, 109.72 g CO<sub>2</sub> eq m<sup>-2</sup> yr<sup>-1</sup>. The highest CH<sub>4</sub> areal flux occurred in PRB (300.45 g CO<sub>2</sub> eq m<sup>-2</sup> yr<sup>-1</sup>), while the lowest in YeRB (54.97g CO<sub>2</sub> eq m<sup>-2</sup> yr<sup>-1</sup>). PRB had the highest N<sub>2</sub>O areal flux of 4059.25 g CO<sub>2</sub> eq m<sup>-2</sup> yr<sup>-1</sup>, while SeB had the lowest, 1086.28 g CO<sub>2</sub> eq m<sup>-2</sup> yr<sup>-1</sup>.

Fig. 3b presents that the hydropower reservoir CI varied largely across the China's hydropower reservoirs. Reservoirs in high-altitude regions (>1000 m) generally had lower CIs (average =  $18.09 \text{ kg CO}_2$ eq MWh<sup>-1</sup>), such as those in YRB (13.92) and SwB (15.39), while lowaltitude reservoirs had higher CIs (average =  $156.39 \text{ kg CO}_2 \text{ eq MWh}^-$ <sup>1</sup>). However, there are several abnormal cases. For example, the value of the Longyangxia hydropower reservoir (altitude = 2582 m) is considerably higher than the average in high-altitude regions, mainly because of its larger area compared to other hydropower stations in high-altitude areas. In general, reservoirs in most high-altitude areas tend to have smaller surface areas. However, the value of the Longyangxia hydropower reservoir is more than 10 times the average in high-altitude regions. The Longyangxia reservoir had a larger surface area primarily as it is the first large-scale cascaded hydropower station on the mainstream of the Yellow River. It has a relatively large installed capacity of 1280 MW. The Shuikou hydropower station is located in a low-altitude area (altitude = 63 m), but its CI (49.5 kg  $CO_2$  MWh<sup>-1</sup>) is considerably lower than the average in low-altitude reservoirs. The Shuikou Hydropower



Fig. 2. Emissions intensity of hydropower reservoir. The x-axis represents the results of GHG emissions on a logarithmic scale with a base of 10. (a) areal flux; (b) carbon intensity.



Fig. 3. Emissions intensity from the studied reservoirs at basin scale. (a) GHG areal flux (b) carbon intensity.

reservoir had a lower CI because its installed capacity is 44.10 % larger than the average at the same altitude.

The International Energy Agency (IEA) indicates that a CI from a reservoir lower than 80 kg  $CO_2$  eq MWh<sup>-1</sup> can be regarded to correspond to a sustainable electricity generation (IHA, 2010). This value is consistent with the energy-related goals of the United Nations 2030 Agenda for Sustainable Development. In this study, 70 reservoirs meet this target, while the average CI of the remaining nine reservoirs is 406.78 kg  $CO_2$  eq MWh<sup>-1</sup>. This highlights the mitigation potential for those nine reservoirs. In their total emission, N<sub>2</sub>O, CO<sub>2</sub>, and CH<sub>4</sub> contributed with 86.03 %, 9.55 %, and 4.42 %, respectively. For the other 70 hydroelectric power stations, N<sub>2</sub>O, CO<sub>2</sub>, and CH<sub>4</sub> contributed

with 76.35 %, 16.71 %, and 6.94 % in the total GHG emission, respectively. The proportion of N<sub>2</sub>O emissions from those nine hydropower reservoirs is larger. Furthermore, the GWP of N<sub>2</sub>O is 298, considerably larger than those of the other two gases. Therefore, the nine hydropower stations that exceed the emission standards need to implement further measures to reduce N<sub>2</sub>O emissions.

# 3.2. Relationship between the reservoirs' GHG emissions and hydropower indicators

As shown in Fig. 4, the installed capacity of hydropower was significantly positively correlated with all GHG pathways except the CH<sub>4</sub>



Fig. 4. Heatmap of relationship among GHG emissions and hydropower reservoir indicators. Pearson correlation analysis was employed.

bubbling emissions. This is consistent with the current understanding that the CH<sub>4</sub> bubbling emissions are from CH<sub>4</sub> produced through anaerobic digestion in sediments, which is mainly affected by temperature and hydrostatic pressure (Hertwich, 2013). The CO<sub>2</sub> diffusive emissions exhibit highly significant positive correlations with the reservoir area, reservoir capacity, and Water residence time (WRT) ( $P \leq$ 0.005) and weak positive correlation with the precipitation ( $P \le 0.05$ ). The CH<sub>4</sub> diffusive emissions exhibit highly significant positive correlations with the reservoir area, reservoir capacity, and WRT. The CH<sub>4</sub> bubbling emissions have highly significant positive correlations with the reservoir area, reservoir capacity, and WRT and weak positive correlations with the air temperature and horizontal radiation. The CH4 degassing emissions exhibit highly significant positive correlations with the reservoir area, reservoir capacity, and WRT, significant positive correlations with the precipitation and air temperature (P < 0.01), and weak positive correlation with the hydropower station-controlled catchment area. The total CH<sub>4</sub> emissions exhibit highly significant positive correlations with the reservoir area, reservoir capacity, and

WRT and weak positive correlations with the precipitation, air temperature, and hydropower station-controlled catchment area. The N<sub>2</sub>O emissions have highly significant positive correlations with the reservoir area, reservoir capacity, reservoir NH<sup>4</sup><sub>4</sub> concentration, age, and WRT, significant positive correlation with the reservoir NO<sub>3</sub> concentration, and weak positive correlation with the total nitrogen concentration. The total GHG emissions exhibit highly significant positive correlations with the reservoir area, reservoir capacity, reservoir NH<sup>4</sup><sub>4</sub> concentration, and WRT, which suggests that these four indicators have a key role in the GHG emissions from the studied reservoirs.

#### 3.3. Projected GHG emissions from reservoirs

As shown in Fig. 5, 1959 sample data related to temperature, precipitation, and river flow (accounting for 80 % of the dataset) were selected as a training set, while the remaining 490 samples (accounting for 20 % of the dataset) were selected as a test set for model training and validation. The results indicated a high degree of accuracy with  $R^2$  =



Fig. 5. Simulations of streamflow for hydropower reservoirs by RF model. (a) and (b) present the results of training and validation processes, respectively.

0.97 for the training set. Similarly, the test set had a satisfactory performance, with  $R^2 = 0.94$ . Thus, the prediction under future climate changes through RF is reliable.

In 2020, the GHG emissions from the studied reservoirs were 17.66 Tg CO<sub>2</sub> eq yr<sup>-1</sup>. As shown in Fig. 6a, under the SSP119 scenario in 2025, the GHG emissions decrease to 17.07 Tg  $CO_2$  eq yr<sup>-1</sup>, increase to 17.85 Tg CO<sub>2</sub> eq yr<sup>-1</sup> in 2030, decrease to 17.07 Tg CO<sub>2</sub> eq yr<sup>-1</sup> in 2035, rebound to 18.21 Tg  $CO_2$  eq yr<sup>-1</sup> in 2040, and slowly decrease to 18.04 Tg CO<sub>2</sub> eq yr<sup>-1</sup> in 2045. By 2050, the emissions reach a peak of 18.71 Tg  $CO_2$  eq yr<sup>-1</sup>, followed by a gradual decrease over the next decade to 17.89 Tg CO<sub>2</sub> eq yr<sup>-1</sup>. In contrast, as shown in Fig. 6b, under SSP245 in 2025, the GHG emissions decrease to 17.09 Tg  $CO_2$  eq yr<sup>-1</sup>, increase to 17.25 Tg CO<sub>2</sub> eq yr<sup>-1</sup> in 2030, decrease again to 17.16 Tg CO<sub>2</sub> eq yr<sup>-1</sup> in 2035, rebound to 18.00 Tg  $CO_2$  eq yr<sup>-1</sup> in 2040, and decrease to 17.18 Tg  $CO_2$  eq vr<sup>-1</sup> in 2045. Over the following 15 years, the emissions gradually increase, reaching a peak of 18.83 Tg CO<sub>2</sub> eq yr<sup>-1</sup> in 2060. As shown in Fig. 6c, under SSP585, the GHG emissions decrease to 17.37 Tg CO<sub>2</sub> eq yr<sup>-1</sup> in 2025, and then gradually increase. In 2060, the emissions reach a peak of 21.02 Tg  $CO_2$  eq yr<sup>-1</sup>

The variation in GHG emissions across basins is notable under the studied scenarios. Under the SSP119 scenario, the emissions of the SeB, YRB, and PRB reservoirs fluctuate for the first 30 years, and ultimately decrease but still exceed the current emission levels (Fig. 6a). On the other hand, the emissions from the reservoirs in SLRB, YeRB, and SwB decrease in the period of 2020 to 2035, followed by an increase in the period of 2035 to 2040, and stabilize by 2060. Notably, by 2060, the reservoirs in SLRB and YeRB are predicted to have lower GHG emissions than their current levels. Conversely, the reservoirs in YRB and PRB are expected to emit slightly more GHG emissions than the current level, while the reservoirs in SwB will maintain their current emission levels. In Fig. 6b, under the SSP245 scenario, the SeB, SLRB, YeRB, SwB, YRB, and PRB reservoirs exhibit fluctuations for the first 40-year period. By 2060, only SLRB is expected to have lower GHG emissions than the current levels. Finally, under the SSP585 scenario (Fig. 6c), the emissions from the reservoirs in YRB and PRB experience slight decreases in the period of 2020 to 2025, followed by sharp increases. The reservoirs in SeB, SLRB, YeRB, and SwB fluctuate over the next 20 years, and

gradually surpass the 2020 emission levels.

According to GHG category, Fig. 6d shows that, by 2060, CH<sub>4</sub> and CO<sub>2</sub> increase by 19.97 % and 17.14 %, respectively, under the SSP119 scenario. However, the N<sub>2</sub>O emission fluctuates and decreases to 13.95 Tg CO<sub>2</sub> eq yr<sup>-1</sup> in 2060. Fig. 6e shows that, compared to 2020, the N<sub>2</sub>O and CH<sub>4</sub> emissions in 2060 are expected to increase by 4.76 % and 51.08 % under the SSP245 scenario, respectively, while the CO<sub>2</sub> emissions are predicted to decrease by 1.17 %. Fig. 6f illustrates a significant increase in GHG emissions, among which the N<sub>2</sub>O, CH<sub>4</sub>, and CO<sub>2</sub> emissions are predicted to increase by 11.02 %, 81.10 %, and 41.34 % in 2060 under the SSP585 scenario, respectively.

#### 4. Discussion

## 4.1. Environmental factors and their effects on GHG emissions from reservoirs

Our results show that N<sub>2</sub>O contributed with more than 80 % of the total GHG emissions in 2020. The above correlation analysis shows that the NH<sup>+</sup><sub>4</sub> concentration in the reservoir mainly from agriculture and industry is an important factor affecting the N<sub>2</sub>O emissions from hydropower reservoirs. The first step of nitrification is to oxidize NH<sup>+</sup><sub>4</sub> to NH<sub>2</sub>OH and generate N<sub>2</sub>O and NO<sub>2</sub> simultaneously. NH<sub>2</sub>OH is then further oxidized to N<sub>2</sub>O or NO<sub>2</sub> is reduced to N<sub>2</sub>O. In other words, NH<sup>+</sup><sub>4</sub> promotes N<sub>2</sub>O production in the reservoir when a sufficient content of NH<sup>+</sup><sub>4</sub> is available (Cheng et al., 2019; Liu et al., 2017). This indicates potential N<sub>2</sub>O mitigation strategies by reducing the NH<sup>+</sup><sub>4</sub> concentration. Besides, high nitrogen loads due to agricultural runoff can also significantly increase the N<sub>2</sub>O emissions (Huang et al., 2024).

Notably, we did not observe significant variations in  $CO_2$  emissions and  $CH_4$  emissions with the reservoir age, compared to other studies (Barros et al., 2011). However, we revealed a significant correlation between the reservoir age and  $N_2O$  emissions, which has been seldom reported previously. As the reservoir ages, the river carries more sediment into the reservoir. The nitrification and denitrification of sediments are strengthened and result in an increase in  $N_2O$  emissions.

Further, the water residence time is an important factor affecting the



Fig. 6. Projected GHG emissions from reservoir categorized by basin (a, b, c) and GHG category (d, e, f).

GHG from all pathways, which is consistent with the report by IHA (2010). When the water residence time is longer, the organic matter in the water body has more time to decompose. These organic matters, under the action of microorganisms, are converted into CO<sub>2</sub>, CH<sub>4</sub>, and other products through various biochemical reactions. Therefore, a longer water residence time promotes the emission of CO<sub>2</sub> and CH<sub>4</sub>. N<sub>2</sub>O is also one of the products of further transformation of NO<sub>2</sub>, an intermediate product in the nitrification process. The extension of hydraulic retention time may provide more time for nitrifying bacteria to oxidize NH<sup>4</sup><sub>4</sub> into NO<sub>2</sub>, and thereby increase the potential for N<sub>2</sub>O generation. N<sub>2</sub>O can also be produced during the denitrification process, particularly when denitrification conditions are incomplete (presence of dissolved oxygen, insufficient C/N ratio, etc.).

The pivotal role of air temperature in influencing GHG emissions in reservoirs has been reported (DelSontro et al., 2010; Kankaala et al., 2004; Xing et al., 2005; Yang et al., 2014). An increase in air temperature can expedite the decomposition of organic carbon, and thereby foster higher GHG emissions. Furthermore, the optimal water temperature for CH<sub>4</sub> production is approximately 25 °C (Dunfield et al., 1993), while, for denitrification and nitrification reactions in water, the optimal water temperatures are 25 and 23 °C, respectively (Garnier, J., 2000). For benthic nitrification and denitrification reactions, the optimal water temperature is 20 °C (Billen et al., 2015). Therefore, in the range where bacteria can tolerate fluctuations, the production of CH<sub>4</sub> and N<sub>2</sub>O increases with the water temperature. Owing to the highest water temperature tolerance of methane-producing bacteria, this can also explain the highest increase in methane in future climate changes.

The GHG emissions from reservoirs align with the overarching trend of climate change. The comparison of the present weather state to future scenarios reveals a significant increase in the annual average air temperature. Under the SSP585 scenario, the annual average air temperature has the largest increase (3.55 °C), and, accordingly, the GHG emissions from the reservoir increase most significantly (an increase of 3.36 Tg CO<sub>2</sub> eq yr<sup>-1</sup>). In addition, among various future climate scenarios, the GHG emission growth rate of the SwB reservoir is lowest (0.34 %, 1.20 %, and 1.34 %), likely as the air temperature increase in SwB is slowest.

#### 4.2. Comparison to previous studies

There are numerous studies on evaluations of GHGs from hydropower reservoirs, which generally ignored the N<sub>2</sub>O emissions (Li and He, 2022; Li et al., 2015). Our study fills this gap by considering N<sub>2</sub>O. Table 2 compares previous estimates to our results, which can be divided into model-based and measurement-based results. The results varied due to different methods and data sets of estimation. Lauerwald et al. (2019) estimated that the N<sub>2</sub>O emission flux from the global reservoirs is 62.94 g CO<sub>2</sub> eq m<sup>-2</sup> yr<sup>-1</sup> based on the method of Maavara et al. (2019). This estimate is considerably lower than our estimate of 1665.8 g CO<sub>2</sub> eq m<sup>-2</sup>

Table	2
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Comparison of results with earlier studies Estimates of GHG emissions from hydropower reservoirs.

vr<sup>-1</sup>, as they did not consider the nitrification and denitrification of sediment, which contributed with almost half of the total N2O emissions in this study. Most notable are the reservoirs of Baishan, Guanyinyan, and Lubuge, which account for more than 65 % of the total N<sub>2</sub>O emissions. The process of generation of N2O in reservoir sediments mainly involves microbial nitrification and denitrification, whereby N2O is produced under aerobic conditions by nitrifying bacteria in the process of converting NH<sup>+</sup><sub>4</sub> to NO<sub>3</sub> (He et al., 2017). Numerous studies have also shown that nitrification and denitrification in sediments are important sources of N<sub>2</sub>O emissions (Deutsch et al., 2010; Usui et al., 2001). The N<sub>2</sub>O emissions evaluated by Li et al. (2018) and Deemer et al. (2016) by measurement-based approaches are considerably lower than ours, which can be explained as many commonly used techniques for measurement of aquatic GHG emissions focus on quantification of the diffusion flux of gases at the air-water interface. For N<sub>2</sub>O, the primary flux pathway is based on the high solubility in water (with a molar fraction solubility of 5.07  $\times$   $10^{-4}$  at 20 °C). Consequently, the  $N_2O$ emissions obtained through these measurement-based approaches are significantly lower than the actual N<sub>2</sub>O emissions from reservoirs, as measurement-based approaches cannot include the whole emissions of reservoirs by setting sampling points. For instance, when hydroelectric power plants are dredged, the churning of large amounts of sediment in reservoirs also increases the N<sub>2</sub>O emissions, which may be ignored by the measurement-based approaches (Cheng et al., 2019).

Our simulation results for CO<sub>2</sub> emission from reservoirs are close to those of Deemer et al. (2016) and Hertwich (2013) (Table 2). This supports the applicability of our G-res model-based approach. Li et al. (2015) employed a hybrid method by combining measurement-based and model-based approaches to predict CO2 and CH4 emissions from a hydropower reservoir. They estimated that the CO<sub>2</sub> emissions are roughly twice higher than those in our study, as they did not consider the emission before the reservoir construction, which was considered with the G-res tool. Li et al. (2018) estimated that the CO<sub>2</sub> emissions from China reservoirs are also twice higher than ours, as most of their reservoirs were in low-altitude areas. Most of the reservoirs in this study were located in high-altitude areas, with 2-3 freezing months annually. Freezing effects can significantly decrease GHG emissions from reservoirs (Boereboom et al., 2012). Under the freezing effect, the ice on the surface of the reservoir blocks the gas exchange at the water-gas interface, which includes the exchange of oxygen with the reservoir water as well as the emission of CO<sub>2</sub>. In addition, the freezing effect reduces the microbial activity, which, in turn, reduces the CO<sub>2</sub> emission.

The CH<sub>4</sub> emissions estimated by Li et al. (2015) are half of ours, mainly because they did not consider CH<sub>4</sub> degassing emissions. Our study demonstrates that CH<sub>4</sub> degassing emissions contributed with 52.50 % of the total CH<sub>4</sub> emissions. The CH<sub>4</sub> diffusion emissions and CH<sub>4</sub> bubbling emissions estimated in this study (62.95 g CO<sub>2</sub> eq m<sup>-2</sup> yr<sup>-1</sup>) are close to those of Li et al. (2015). A detailed description on CH4 degassing has been reported by Hertwich (2013).

<b>T</b>									
	Area (10 <sup>6</sup> km <sup>2</sup> )	Areal flux		Method	Source				
		CO <sub>2</sub>	CH <sub>4</sub>	N <sub>2</sub> O					
		g CO <sub>2</sub> eq m <sup>-2</sup> y	r <sup>-1</sup>						
China's 79 reservoirs	0.02	362.99	132.54	1665.8	Model	This study			
China reservoirs	0.03	712.65	169.60	173.66	Measurement	Li et al., 2018			
Global reservoirs	0.25			62.94	Model	Lauerwald et al., 2019			
Global reservoirs	0.31	441.65	1985.6	51.28	Measurement	Deemer et al., 2016			
China reservoirs	0.03	867.24	65.52		Measurement	Li et al., 2015			
					and model				
Boreal reservoirs	0.08	970	1360		Model	Hertwich, 2013			
Temperate reservoirs	0.13	420	288		Model	Hertwich, 2013			
Tropical reservoirs	0.12	1200	1840		Model	Hertwich, 2013			
Temperate reservoirs	0.9	511	248.2		Measurement	St. Louis et al., 2000			
Tropical reservoirs	0.6	1277.5	3723		Measurement	St. Louis et al., 2000			

#### 4.3. Uncertainty analysis

Common sensitivity analysis methods include local sensitivity analysis and global sensitivity analysis (Borgonovo and Plischke, 2016). Global sensitivity analysis is the recommended method in the field of hydrology (Pappenberger et al., 2008), because, compared to local methods, global methods offer specific advantages, including the ability to simultaneously consider the impact of changes in each parameter on the model output across the entire parameter range, and high degree of adaptability to nonlinear and nonmonotonic models. Therefore, we employ a global sensitivity analysis to verify the uncertainty of the results, with the widely used Sobol' method (Ganji et al., 2016).

Parameters involved in GHG emission calculations, obtained from measurements or modeling (Fig. 7), can increase the degree of uncertainty. Therefore, the sensitivity of these parameters was tested based on a Sobol sensitivity analysis with 4000 sample points.

As shown in Fig. 7a, the effect index of the reservoir soil carbon content (0.74) on the CO<sub>2</sub> emissions was largest, followed by those of the reservoir area (0.29), river area before impoundment (0.25), air temperature (0.19), and total phosphorus (0.01). The air temperature has the largest effect index for CH<sub>4</sub> emissions (Fig. 7b), accounting for 48.43 % of the total effect index, followed by the littoral area (18.12 %), WRT (13.99 %), and catchment area (11.69 %), while the other factors together account for only 7.77 %. The water temperature, reservoir area, and NH<sup>+</sup><sub>4</sub> were the top three factors for N<sub>2</sub>O emissions (Fig. 7c), which contributed with 89.43 % of the total effect index.

Based on the sensitivity analysis results, the confidence interval was explored by a Monte Carlo simulation under a change of  $\pm 30$  %. The results of 10,000 samples are normally distributed; the highest bar represents the highest frequency (Fig. 8). To quantify the uncertainty, we utilized a confidence interval of 5 % to 95 %.

As shown in Fig. 8a, the 95 % confidence interval of CO<sub>2</sub> annual emissions ranged from  $2.02 \times 10^7$  to  $4.05 \times 10^7$  kg CO<sub>2</sub> yr<sup>-1</sup>. The highest-frequency value  $(2.92 \times 10^7$  kg CO<sub>2</sub> yr<sup>-1</sup>) is close to our calculated result (mean value:  $2.95 \times 10^7$  kg CO<sub>2</sub> yr<sup>-1</sup>). In Fig. 8b, the 95 % confidence interval of CH<sub>4</sub> annual emissions ranges from  $9.67 \times 10^6$  to  $1.61 \times 10^7$  kg CO<sub>2</sub> eq yr<sup>-1</sup>. The highest-frequency value of  $1.24 \times 10^7$  kg CO<sub>2</sub> eq yr<sup>-1</sup>). Fig. 8c shows that the 95 % confidence interval of N<sub>2</sub>O annual

emissions ranges from  $1.03\times10^8$  to  $2.51\times10^8$  kg CO<sub>2</sub> eq yr<sup>-1</sup>. The highest-frequency value (1.75  $\times10^8$  kg CO<sub>2</sub> eq yr<sup>-1</sup>) is close to our calculated result of  $1.81\times10^8$  kg CO<sub>2</sub> eq yr<sup>-1</sup>. In summary, the Sobol sensitivity analysis and Monte Carlo simulations demonstrate the reliability of our results.

#### 4.4. Mitigation of GHG emissions from hydropower reservoirs

Although previous studies generally indicated that the GHG emission intensity of hydropower is considerably lower than that of thermal power generation (Feng et al., 2014; Hondo, 2005), our results demonstrate that the considerable GHG emissions from hydropower reservoirs, ignored in previous studies, should be carefully considered, as they can hinder the climate mitigation goal, particularly as China will continue to develop the hydropower according to the 14th Five-Year Plan. By 2030, the installed capacity of hydropower is expected to reach approximately 370 million kW. Below, we suggest strategies to reduce the GHG emissions from reservoirs.

For existing hydropower reservoirs, targeted measures should be employed considering their specific emission pathways and characteristics. Our results show that the N<sub>2</sub>O dominates in the GHG emissions, particularly for the nine reservoirs that cannot meet the IEA's target for a sustainable hydropower. In addition, regular maintenance of equipment and dredging of reservoirs are necessary. Notably, China has already begun to consider the issue of sediment deposition in older hydropower reservoirs. In 2021, sediment deposition was enacted for a large hydropower station on the Yellow River (Zhou et al., 2021). However, the sediment is still the main factor driving GHG emissions in the studied reservoirs, which requires more specific solutions. For instance, the N<sub>2</sub>O emissions from sediments in the Baishan, Xinanjiang, and Lubuge reservoirs account for more than 65 % of the N<sub>2</sub>O emissions. Measures such as scouring of the lower riverbed to expel and reduce the sediment should be implemented for these reservoirs.

Finally, considering the significant positive correlation between  $NH_4^+$ and  $NO_3^-$  with  $N_2O$  emissions from reservoirs, measures need to be implemented to improve the water quality and mitigate the eutrophication (for example, by reducing artificial nutrient inputs and applying sludge removal treatments). Reservoirs with high  $NO_3^-$  and  $NH_4^+$  concentrations, such as those in Xiaolangdi Reservoir, Fengman Reservoir,



Fig. 7. The Sobol' total effect indexes of GHG emission. (a) CO<sub>2</sub>;(b) CH<sub>4</sub>;(c) N<sub>2</sub>O.



Fig. 8. The uncertainty analysis results of GHG emissions. (a) CO<sub>2</sub>;(b) CH<sub>4</sub>;(c) N<sub>2</sub>O.

and Tingzikou Reservoir (Fig. S2), are generally located in China's major grain-producing areas. However, these areas are still exposed to an excessive application of nitrogen fertilizers, which can result in higher concentrations of  $NO_3$  and  $NH_4^+$ . Although China implemented measures, including organic and chemical fertilizers, straw recycling, and deep placement of fertilizers, to reduce the agricultural non-point-source pollution (Duan et al., 2024), the pollution in flood areas such as reservoirs should be carefully considered.

For the future hydropower development, our study demonstrates that reservoirs in high-altitude areas generally have lower CIs, and that the steeper terrain in high-elevation areas can favor hydropower projects with higher power densities (Almeida et al., 2019). This provides insightful implications for the future hydropower planning with respect to GHG emission reduction. This also requires to consider other important issues such as biodiversity, ecosystem service, and geological factors. In the design and operation of reservoirs, it is essential to fully consider the potential impact of hydraulic retention time on GHG emissions and employ corresponding measures to reduce these emissions. For instance, optimization of reservoir dispatching, increase in water body mobility, and improvement in water quality can effectively mitigate the adverse effects of long hydraulic retention times on GHG emissions.

#### 4.5. Limitations and prospect of this study

As for previous studies conducted at national and global scales, our estimates of GHG emissions from China's hydropower reservoirs are inevitably impacted by data limitations and uncertainties. In this study, we did not encompass all hydropower reservoirs in China primarily due to data deficiencies, which potentially hinders a comprehensive understanding of GHG emissions from hydropower reservoirs. However, more than 60 % of the country's hydropower generation across all basins was included in this study. Our study also ignores the impact of future changes in air temperature and precipitation on the water quality, which may lead to uncertainty of our projection to some degree. The prediction of changes in TP, NH<sub>4</sub><sup>+</sup>, organic nitrogen, dissolved oxygen, and other factors, mainly resulting from the social and economic development, under the climate changes requires involvement of multiple complex models, in which large uncertainties can emerge. The accurate prediction of socioeconomic development is beyond the scope of this study. Additionally, our projections did not cover all SSP scenarios, though the most representative scenarios were considered. Finally, our data on NO3 and N<sub>2</sub>O in reservoirs were sourced from the literature rather than from measurements, which may increase the uncertainty associated with our findings. However, the uncertainty analysis demonstrated the reliability of our results.

As subjects of future researches, the establishment or enhancement of reservoir-related monitoring systems can include more hydropower reservoirs and collect more specific data, which can provide a more holistic understanding of the impact of climate changes on GHG emissions from reservoirs and related flooded land. In addition, the development of interdisciplinary models integrating climate, energy supply and demand, geoscience, and socioeconomic system can provide more accurate and region-specific results. Finally, a broader range of climate scenarios should be considered in the future to elucidate more detailed emission pathways and provide related mitigation strategies.

#### 5. Conclusion

Mitigation of GHG emissions from hydropower reservoirs has been

widely recognized, especially given that China's ambitious goal to increase its hydropower share. Uncovering and projecting the spatialtemporal pattern of these emissions can therefore prepare insightful implications towards sustainable hydropower. However, previous studies seldomly consider future climate impacts on these emissions. This study fills this gap by combining G-res tool, IMAGE-DGNM and random forest model based on SSP scenarios. 79 hydropower, contributing more than 60 % of the nation's total hydropower generation and across main China's basins, were taken as study cases. We observed significant spatial heterogeneity of GHG emission intensity, in which reservoirs with lower CI were generally located in high altitude area. N<sub>2</sub>O dominates these GHG emissions, followed by CO<sub>2</sub> and CH<sub>4</sub>. The projection further suggests that these emissions will increase under all the three scenarios, in which CH4 has the greatest increase due largely to it is more sensitive to air temperature. Finally, this study provides insightful suggestions to reduce these emissions that have global implications.

#### CRediT authorship contribution statement

Hongqiao Chen: Writing – original draft, Methodology, Investigation. Hengyu Pan: Writing – review & editing, Writing – original draft, Supervision, Methodology, Funding acquisition, Data curation, Conceptualization. Shijiang Xiao: Writing – review & editing, Validation, Methodology. Shihuai Deng: Supervision.

#### Declaration of competing interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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#### Supplementary materials

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#### Data availability

Data will be made available on request.

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